Appraisal of geochemical composition and hydrodynamic sorting of the river suspended material: Application of time-integrated suspended sediment sampler in a medium-sized river (the Sava River catchment)

Mavro Lučić<sup>a</sup>, Nevenka Mikac<sup>a</sup>, Niko Bačić<sup>a</sup>, Neda Vdović<sup>a\*</sup>

<sup>a</sup> Division for Marine and Environmental Research, Ruđer Bošković Institute, Bijenička cesta
54, 10000 Zagreb, Croatia

Mavro.Lucic@irb.hr, Nevenka.Mikac@irb.hr, Niko.Bacic@irb.hr

\*corresponding author, Neda.Vdovic@irb.hr

## ABSTRACT

1 The suspended particulate matter (SPM) carried by the rivers shows a wide range of particle 2 size classes, mineralogical and chemical compositions and is mainly influenced by 3 hydrodynamic sorting and provenance during the transport. Here, we have investigated the 4 composition of the SPM in the Sava River and its tributaries (Ljubljanica, Savinja and 5 Krapina) using a time-integrated suspended sediment sampler (TIMS). The 6 representativeness of material collected by TIMS was evaluated comparing fine-grained 7 sediments, single-point SPM and SPM collected using a shallow and deep-positioned 8 sampler. The main results have revealed that the mineralogical and geochemical 9 composition of the material is largely dependent on hydrological conditions. The 10 differentiation of element composition is especially emphasized at low water stage when 11 most of the SPM consists of slow-settling mineral phases (clay minerals and metal 12 oxyhydroxides) which can be trapped in the sampler. During periods of high discharges, 13 differentiation is less prominent, and homogenization of the SPM occurs, mainly as a part of

14 bed load is also taken into suspension. These conditions have proven unfavorable for 15 sampler efficiency, as at least part of the finest particles could not be retained. Additional 16 issues that may occur during TIMS employment relate to biologically driven carbonate 17 precipitation, which is triggered by changes in physico-chemical conditions at low water table 18 in the summer period. Increased concentration of Ca, related to that process, influences the 19 elemental composition of the SPM, which is particularly important when anthropogenic 20 impact or sediment source is assessed. Hence, in order to interpret the geochemical and 21 mineralogical data collected by TIMS, these factors should be taken into account. Our 22 findings emphasize the need for detailed studies of chemical composition of the SPM (time-23 integrated) in medium-sized rivers and point out the significance of evaluating sampling 24 representativeness during different hydrological conditions.

Keywords: Geochemical composition; Hydrodynamic sorting; Time-integrated sampler; the
 Sava River catchment

### 27 **1. Introduction**

28 The suspended particulate material (SPM) refers to particles that suspend in the water 29 column with a lower size limit of 0.20 or 0.45 µm in median diameter (Viers et al. 2009). It 30 consists of inorganic (quartz, feldspars, carbonates, clay minerals, metal oxyhydroxides, 31 heavy minerals) and organic (microorganisms and detritus) particulate matter (Gregory, 2006; Garzanti et al., 2011) usually present in flocculated form (Droppo and Ongley, 1994). 32 33 The SPM has a major role in transfer of elements from source to sink. According to Horowitz 34 (1991) and Audry et al. (2004), more than 90% of the riverine flux of metals is associated 35 with fine-grained sediment. So, when dealing with the trace elements input/transfer/transport 36 along a river, the investigation should focus on the SPM. However, most of the SPM-bound 37 element load is related to high flow events which are extremely irregular. Another difficulty is 38 to collect a representative sample and sufficient amount of the SPM for analyzing different 39 chemical and/or physical characteristics. In order to meet these requirements, Phillips et al. 40 (2000) developed a time-integrated suspended sediment sampler (TIMS), mainly designed

41 for streams and small lowland rivers (Schindler Wildhaber et al. 2012; Smith and Owens, 42 2014). Heretofore, TIMS was used in different studies (Russel et al., 2000; McDonald et al., 43 2010; Droppo et al., 2019) and proven effective in several environmental and controlled 44 laboratory conditions (Martínez-Carreras et al., 2012; Marttila et al., 2013; Perks et al., 2014). 45 The amount of material collected was found sufficient and the representativeness of TIMS 46 was also proven satisfying. The main objections were the poor assessment of the SPM mass 47 flux (Goharrokhi et al., 2019) and lack of knowledge of how it operates in larger river systems 48 (Smith and Owens, 2014).

TIMS was also used in our previous research (Lučić et al., 2019) to investigate the Sava River SPM and associated anthropogenic impact. The TIMS was set at one location in the Sava River in Zagreb during different discharge periods. The results have shown increased concentrations of some ecotoxic elements (As, Bi, Cd, Cr, Ni, Pb, Sb, Zn). Some issues opened during that investigation: the input of the material from different sources during different discharge periods and high concentration of calcite in spring sampling period was observed. A possibility of *in situ* calcite precipitation instigated by algal bloom was

56 hypothesized.

57 This investigation has been conducted in order to obtain more information on the

58 hydrodynamic sorting and representativeness of the SPM (time-integrated) transported in a

59 medium size river. The case study was the Sava River and main tributaries (Ljubljanica,

60 Savinja and Krapina) between its source and the city of Zagreb, as an example of medium

- 61 size river in the anthropogenically impacted environment. The aims of the study were:
- to characterize the spatial and time variation of the geochemical composition of the
   SPM sampled by TIMS in different hydrological conditions,
- 64 2. to assess potential influence of hydrodynamic sorting on suspended material,
- 65 3. to determine possible differences between shallow and deeper suspended load using
   66 two TIMS samplers at one location.
- 67
- 68

#### 69 **2. Materials & methods**

#### 70 2.1. Study area

The Sava River is a major Danube tributary which flows through Slovenia and Croatia, alongside the northern border of Bosnia and Herzegovina, and finally through Serbia. The upper course of river is 232 km long with drainage basin covering 12680 km<sup>2</sup> of surface area (Table 1). Other detailed geographical characteristics of the Sava River can be found in our previous work (Lučić et al., 2019).

76 The Sava River has its origin as two branches, the Sava Dolinka and Sava Bohinjka rivers 77 which flow mainly through carbonate terrain (limestone and dolostone) until their confluence 78 at the city of Radovljica. From there, the river passes through previously deposited fluvio-79 glacial terraces which alternate with Paleozoic rocks consisting of shales, quartz sandstones 80 and conglomerates. Besides Permo-Carbonian clastic sediments, the central part of the Sava 81 section in Slovenia is composed of Triassic carbonates, together with Paleogene and 82 Neogene clastic rocks in the area before the border with Croatia (Placer, 2008). In Croatia, 83 the drainage area of the river comprises terraced Quaternary deposits consisting of sands, 84 marls and clays (Šikić et al., 1979). 85 The Ljubljanica River flows through Ljubljana Moor, paleo-swamp filled with Quaternary 86 alluvial sediments (pebble, sand and clay) covering Paleozoic basement and Mesozoic 87 limestone and dolostone. Because of the considerable thickness of fluvio-glacial sediments, 88 the Ljubljansko Barje is one of the most important aquifers in Slovenia (Cerar and Urbanc,

89 2013).

The Savinja River, the second longest Slovenian river originates in the Kamnik-Savinja Alps
 in Triassic carbonates and flows through Oligocene tuffs and andesites, while in the lower

92 part drains the Triassic carbonates and Miocene sandstones. As a result of its runoff

93 characteristics, its catchment area contributes up to 40 % of the lower Sava River (in

94 Slovenia) discharge in high rainfall events (Kobold and Sušelj, 2005).

95 The Krapina River has its origin in the Paleogene and Neogene of the Panonnian Basin. On

96 the left side of the flow, the Krapina River is filled by many streams that drain the Medvednica

97 mountain, consisting of rocks of different ages (from Silurian to Quaternary age) (Galović and

98 Peh, 2014).

- 99 2.2. Sampling and preparation of samples
- 100 The design and principles of TIMS are thoroughly described in the papers of Philips et al.,
- 101 (2000) and Russel and Walling (2000). It is a PVC cylinder of 100 cm in length and 10 cm
- 102 diameter, closed on one side with a simple screw cap and at the other with the conus shaped
- 103 screw cap. At the center of each cap there is a 4 mm opening to allow the passage of water.
- 104 The sampler is positioned parallel to the flow of the river with the conus part pointing opposite
- 105 to the flow direction. Within the main body of the sampler, the flow velocity is reduced by a
- 106 factor > 600, which induces sedimentation of the suspended sediment.
- 107 The sampler is intended to collect river suspended sediment in a prolonged period. The
- 108 resulting composite sample integrates natural variations in the geochemical properties of
- 109 sediment transported during different runoff events. The sample mass collected by the
- 110 sampler is sufficient to satisfy a wide range of laboratory analyses. Samplers used in this
- 111 study were slightly modified we used larger inlet and outlet (6 mm diameter), while tube
- 112 diameter was 11 cm in width.
- 113 Sampling was conducted during a hydrological year with campaigns organized from October
- 114 **2016** to July 2017 at different frequencies (Table 2); at least two sampling campaigns were
- 115 performed at each site. The samplers were positioned at five locations (Fig. 1). At four of the
- 116 chosen locations, TIMSs were fixed by steel uprights near the riverbed since river channels
- 117 were too shallow for two samplers. At Zagreb location, TIMS was tied under the pontoon and
- 118 sunk 30 cm below the water surface. During the spring and summer periods, an additional
- 119 sampler was fixed closer to the river bottom. The intention was to collect and compare
- 120 shallow and deeper suspended load at the same location.
- 121 **Figure 1.**
- 122 Besides suspended material collected by TIMS, in the period of time-integrated sampling at
- 123 the Sava River, single-point samples of suspended material were also taken, as well as the
- bottom river sediments. The frequency of sampling is shown in Table 2.

- 125 The sediments were sieved through a 63 µm sieve, using ambient water, to separate coarse
- 126 sand fraction and approximate the sediment sample to TIMS and single-point samples; the
- 127 fraction <63 µm was kept for analyses. The samples were used wet for the particle size
- 128 analysis; for all the other analyses the material was freeze-dried (FreeZone 2.5; Labconco)
- 129 and ground to fine powder using a ball-mill (Pulverisette 7; Fritsch).
- 130 For the single-point SPM samples, river water was taken in plastic 6 dm<sup>-3</sup> bottles. A portion
- 131 (500 cm<sup>3</sup>) of each sample was used "as is" for the particle size determination; the rest of the
- 132 samples was subsequently filtered in laboratory (0.45 µm cellulose acetate, Sartorius). The
- 133 material remaining on the filter was then dried at 60°C and kept for the element analysis.
- 134 Materials collected by TIMS were transferred into large glass beakers and left to settle; the
- 135 supernatant was then carefully decanted and sedimented material kept for analyses. A
- 136 portion of sedimented material (~2g) was used "as is" for the particle size determination and
- 137 the rest of it was freeze-dried for other analyses.
- 138 2.3. Methods
- 139 Particle size distribution (PSD) was determined using a laser-based particle size analyzer
- 140 (LS 13320; Beckman Coulter Inc.). The PSD was calculated using Mie theory of light
- scattering (optical parameters: refractive index = 1.53; absorption index = 0.1). PSD was
- 142 determined for the samples in their native state as well as for the organic-free samples. To
- 143 remove organic matter, samples were treated with  $15\% H_2O_2$ .
- 144 The mineral composition was identified by X-ray powder diffraction using a Philips X-Pert
- 145 MPD diffractometer (40 kV, 40 mA, range scanned 4–63° 2θ). Bulk composition of 10 TIMS
- and 7 sediment samples was determined. For 6 chosen TIMS samples, the clay fraction (< 2
- 147 µm) was separated by centrifugation and analyzed on oriented slides after being air-dried,
- saturated by ethylene glycol, and heated for 1 h at 400° and 550°C. Proportions of minerals
- 149 were evaluated semi-quantitatively using distinctive peak areas (Moore and Reynolds, 1997;
- 150 Kahle et al., 2002), weighted by Schultz (1964) empirical factors, which represents rough
- 151 estimate of mineral percentages.

- 152 Prior to geochemical analysis, sediment and TIMS samples were digested in the Microwave
- 153 digestion system Multiwave 3000 (Anton Paar) by a two-step procedure:
- 154 1) 5 cm<sup>3</sup> HNO<sub>3</sub> (65%, pro analysi, Kemika) + 1 cm<sup>3</sup> HCI (37 %, VLSI Grade, Rotipuran)
- 155 **+ 1 cm<sup>3</sup> HF (47-51%, supra pur, Fluka);**
- 156 **2)** 6 cm<sup>3</sup> H<sub>3</sub>BO<sub>3</sub> (≥99.5%, Sigma-Aldrich).
- 157 Due to small amount of sample, single-point SPM samples were dissolved by modified
- 158 procedure:
- 159 1) 4 cm<sup>3</sup> HNO<sub>3</sub> + 1 cm<sup>3</sup> HCl + 0.2 cm<sup>3</sup> HF;
- 160 **2)** 1.25 cm<sup>3</sup> H<sub>3</sub>BO<sub>3</sub>.
- 161 Multi-elemental analysis of dissolved and particulate fraction was conducted using a High-
- 162 Resolution Inductively Coupled Plasma Mass Spectrometer (HR ICPMS), Element 2 (Thermo
- 163 Finnigan). Analytical quality control was provided by simultaneous analysis of blanks and
- 164 certified reference material (Soil-NCS DC 77302) for which good recoveries (90-100 %) were
- obtained, depending on the element measured. Details of the method are provided in Fiket etal. (2017).
- 167 The hydrological data were provided by Meteorological and Hydrological Service of Croatia
- 168 (DHMZ) and Slovenia (ARSO). Discharge measurements were performed by conventional
- 169 current meter method. The SPM concentration was measured daily by filtration of surface
- 170 water samples taken in the middle of the river course, 10 20 cm below the water level.
- 171 Statistical treatments were performed using a R package "robCompositions" while heat-maps
- 172 were designed in package "gplots" in R platform (R Core Team, 2017).

### 173 **3. Results and discussion**

- 174 3.1. Hydrological and particle size characteristics
- 175 Five sampling campaigns conducted in Zagreb encompassed a wide spectrum of discharges
- 176 and concentrations of suspended material (Fig. 2). Except for the calm summer period, all
- 177 seasons were characterized by at least one increase in water discharge followed by
- 178 corresponding variation of the SPM content. Ljubljanica and Savinja rivers have revealed
- 179 similar trends of discharge fluctuations (Fig. A.1 and A.2). Somewhat different hydrological

conditions were observed for the Krapina River in which high SPM concentration did not
always follow the rise of the water level (Fig. A.3), mainly as a result of different sediment
sources (Morehead et al., 2003).

183 **Figure 2**.

- 184 In the Sava River SPM, TIMS and sediment samples, predominant particle-size fraction was
- 185 silt, regardless of the sampling period (Appendix B; Fig. 3). Sediments were dominated by
- 186 finer particle size ranging from 11.1 to 34.7 μm, compared to single-point SPM (19.6 56.3
- $187 \mu$ m) and TIMS samples (40.6 56  $\mu$ m). After organic matter removal, a notable increase in
- 188 clay content and consequently lower mean grain size (Mz) was observed in all samples.
- 189 These changes were more pronounced in TIMS samples than in sediments, which indicates
- 190 that the flocculation process took place inside TIMS. Also, lower Mz in treated sediments
- 191 compared to TIMS samples suggests that the part of the finest material carried in suspension
- 192 could have passed through the sampler in case of a high flow rate. The flocculation of the
- 193 SPM in the river channel could not have been be documented since the material retrieved
- 194 from the single-point SPM sampling was not sufficient to apply H<sub>2</sub>O<sub>2</sub> treatment. However, the
- 195 differences in the PSD of 6 single-point SPM samples retrieved in Sava Zagreb, and
- 196 comparable Mz results between single point SPM and TIMS, imply the presence of
- 197 flocculation at least in some seasons.
- 198 **Figure 3.**
- 199 For the other rivers, the sediment particle size showed comparable variations (Appendix B),
- 200 Mz of the Ljubljanica River ranging from 15.8 to 23.1  $\mu$ m, similar to Savinja (15.4 23.2  $\mu$ m)
- and Krapina rivers (15.1 18.4 µm). Regarding TIMS samples, the ones from the Krapina
- River had lower Mz (19.5 44.2  $\mu$ m), compared to those from Savinja (24.7 71.6  $\mu$ m) and
- 203 Ljubljanica rivers ( $64.9 102.3 \mu m$ ). In all H<sub>2</sub>O<sub>2</sub>-treated samples there was a decrease in Mz.
- 204 The presence of coarser particles in TIMS of the Savinja River was a result of its stronger
- 205 erosive power emphasized during the high-water level when the part of the bed load could
- also be taken into suspension (Singh, 2009). The abundance of sand-sized particles in all

- samples of the Ljubljanica River was largely the effect of erosion of soil aggregates along the
   watercourse (Woodward and Walling, 2007).
- 209 Grain-size data obtained for the shallow (TMZG4) and deep-positioned (TMBZG4) TIMS in
- 210 Zagreb during the spring campaign differ in higher proportion of sand observed in deep-
- 211 positioned sampler. In the summer sampling campaign deep-positioned TIMS (TMBZG5)
- 212 contained higher share of clay fraction. However, when treated samples were compared, Mz
- 213 in shallow positioned TIMSs decreased substantially while no such effect was observed for
- 214 deep-positioned TIMSs. This could be related to the transport of aggregated soil particles in
- shallow relative to deeper load which contained more sand consisting of quartz, carbonates,
- 216 tectosilicates and heavy minerals.
- 217 **3.2.** *Mineralogical characteristics*
- 218 The minerals present in all analyzed samples were quartz, calcite, dolomite, phyllosilicates
- 219 and feldspars (Fig. 4). The most abundant minerals in sediment samples of the Sava River
- were quartz, calcite and dolomite. The dominance of each of them alternated, depending on
- the sampling period. In general, the highest amount of quartz was found in deeper positioned
- samplers and two shallow samplers from the late autumn and spring sampling campaigns.
- 223 Regarding shallow positioned TIMS, higher quartz content probably resulted from stronger
- 224 erosive power during high discharge. In such conditions, material is not sufficiently
- 225 differentiated, mainly because of minimal chemical weathering and the dominance of
- 226 physical erosion in the source area.
- 227 The origin of carbonates in the Sava River SPM and sediments is mostly detrital. Variation of
- dolomite content in TIMS samples can be related to both source supply and hydrodynamic
- sorting; higher density of dolomite affects its accumulation in finer sand and coarser silt
- classes (Garzanti et al., 2009; Garzanti and Ando, 2019). The origin of calcite could be
- 231 double-natured. As assumed for the other minerals, in most of the sampling periods calcite
- was undoubtedly detrital in origin. Nevertheless, the high content of calcite found in sediment
- 233 (55%) and shallow-positioned TIMS (69%) might be related to biologically mediated
- precipitation process (Olivier et al., 2011; Lučić et al., 2019) instigated in the summer period.

235 The higher content of phyllosilicates in TIMS compared to sediments was observed in

236 periods of lower water discharges which suggests that phyllosilicates have preferential

237 transport in surface load and have a good possibility to be retained in the sampler at slow

238 flow rate. Their low content recorded in the SPM and sediment samples of the first and final

sampling campaign was a consequence of dilution effect, mainly controlled by calcite and

240 quartz abundance. Feldspar minerals did not show any variation visible from mineralogical

analyses.

242 Clay mineral composition obtained for TIMS samples of the Sava River was characterized by

243 prevalence of illite/mica minerals, followed by chlorite, smectite, kaolinite and vermiculite,

present only in the summer sampling period (Table A.1). The comparison of clay minerals in

shallow and deeper positioned TIMS, showed a higher amount of illite/mica and the absence

of kaolinite in the latter, which implies that kaolinite does not prefer deeper transport (Gippel,

247 1995).

248 The slightly lower carbonate content was detected in all tributary samples, as the

249 consequence of more siliciclastic lithologies (Fig. 4). All minerals showed opposite behavior

250 relative to quartz. Additionally, TIMS samples had an increased quartz content compared to

sediments. These results were corroborated with grain size analysis which reflected

coarsening of the TIMS samples during high water stages.

253 **Figure 4**.

254 **3.3**. Geochemical composition of suspended material

255 The geochemical composition (50 elements) of the analyzed suspended sediments is given 256 in the Appendix B. The comparison of TIMS samples showed a higher average Ca (107342) 257 mg kg<sup>-1</sup>) and Mg (22761 mg kg<sup>-1</sup>) concentrations in the Sava River than in TIMSs positioned 258 at other locations (Ca – 68310 mg kg<sup>-1</sup> and Mg – 17716 mg kg<sup>-1</sup>); the highest concentrations 259 were observed during the summer sampling period, which corroborated mineralogical 260 records. Magnesium concentrations proved to be sensitive to hydrodynamic sorting with the 261 highest values in TIMSs during the greatest discharge in the Sava (26482 mg kg<sup>-1</sup>) and 262 Savinja (27506 mg kg<sup>-1</sup>) rivers. Other major elements (Al, Fe, Ti, Na, K), being the part of

aluminosilicates, showed the behavior opposite to Ca and Mg, with the highest

264 concentrations in samples of the Krapina River. The rare earth elements (REE) and most of

- the trace elements displayed the highest concentrations in samples with more AI and Fe
- content.

267 **3.4**. Correlation analysis

- 268 Besides absolute element concentrations, geochemical data were also considered in terms
- 269 of their composite nature (Reimann et al., 2012). This means that each variable is part of a
- whole and carry only relative information (Pawlowsky-Glahn et al., 2015). To follow this
- 271 definition, before correlation analysis, we transformed geochemical data using special type of
- 272 log-ratio transformation, called symmetric coordinates (balances), where elements are
- arranged according to a clustering procedure (Kynčlová et al., 2017; Reimann et al., 2017).

274 The correlation analysis (Fig. 5) was conducted separately for all analyzed types of samples;

A) TIMS, B) single-point SPM and C) fine-grained sediments. In these graphs, a similarity

between TIMS and single-point SPM was observed. Fig. 5A (TIMS) is characterized by two

277 large clusters. Starting from the upper-right corner, elements indicate their geogenic nature.

278 Strong clustering is determined for Ga, Sc, V, Ni, Ge, Ti, Li, Al, Rb, K, LREE, Th, HREE, Nb,

279 **Y**, and to a lesser extent for TI, U, Cs and Mg, which are mostly detrital in origin. The second

smaller subcluster consists of Fe, Na, and Cr, the presence of which is assumed in multiple

281 mineral phases (phyllosilicates, feldspars and heavy minerals). Sr and Ca indicate their

incorporation into the carbonate minerals, mostly calcite. The second large cluster of

elements (Co, Ba, Pb, As, Mn, Bi, Be, Cd, Mo, Sn, Sb, Zn, Cu and W) probably emphasizes

their anthropogenic nature and association with the finest particles (Chen et al., 2014).

In a single-point SPM (Fig. 5B) elements such as LREE, Th, Ga, Li, HREE, AI, V, Y, Nb, Ti,

Sc, Cs, Tl, Ge, Rb, Co, Be, Ba, Na, W, Cr and Fe show moderate to strong mutual positive

287 correlation, which suggests their association with clay minerals, oxyhydroxides and organic

- 288 matter. Because most of the single-point samples were taken during low water level and
- strictly in a shallow part of the flow, this group represents geogenic element association
- 290 dominating in a wash load where quartz, tectosilicates, heavy minerals and other fast-settling

phases are sparse. The last cluster consists of two smaller ones. The first one is carbonate
and feldspar related, and consists of U, Mg, Ca, Sr, K and Pb (Garçon et al., 2014); the
second one is dominated by mainly anthropogenic group of elements which tend to
concentrate in the finest part of shallow suspended load, such as Ni, Sn, Bi, Cd, Sb, As, Mn,
Cu, Zn and Mo.

296 The fine-grained sediments (Fig. 5C) display somewhat different clustering and not so 297 pronounced correlations. The upper-right corner suggests agglomeration of elements solely 298 detrital in origin. Moreover, the strong positive correlations of Ti, W, Nb, Ge and Na probably 299 suggest their presence in heavy minerals and tectosilicates which tend to concentrate in 300 deeper suspended load (Garzanti et al., 2011). The second smaller subcluster consists of 301 Ga, LREE, Li and AI mainly associated with clay minerals. An association of elements such 302 as V, Fe, Cr, Sc, As, Pb, Co, Sb, Y and HREE may imply mafic component, especially within 303 sediments of the Savinja and Krapina rivers (Salminen et al., 2005). The heat-map of fine-304 grained sediments does not reveal a prominent anthropogenic association of elements. 305 Rather, they are positioned within the group of elements mostly related to phyllosilicates and 306 feldspars (Rb, Be, Tl, Cu, Bi, Th, K, Ba, Zn, Mn and Cs). The second large cluster consists of 307 two subclusters and point out carbonate components more dominant in the Sava River. 308 Association of elements Cd-Sn is partly influenced by anthropogenic impact. Except for Sn, 309 all grouped elements (Ni, Sr, U, Mg, Cd, Mo and Ca) have a similar ionic radius and can be 310 incorporated into carbonates (Rambeau et al., 2010; Lerouge et al., 2017)

311 **Figure 5**.

# 312 **3.5.** Geochemical normalization

To assess the representativeness of time-integrated sampler for sampling in the Sava River and tributaries, we compared the geochemical composition of the material collected by shallow-positioned TIMS to other sampled material (fine-grained sediment (< 63 µm), deeppositioned TIMS and single-point SPM samples). In this way, important information about geochemical similarity/dissimilarity between material captured in TIMS and other materials present in a water column during different hydrological cycles could be attained. During low

319 discharges, the most of fine-grained sediment is settled to the bottom, and clay/organic 320 particles prevail in suspension. Somewhat different behavior can be observed during the 321 turbulent water conditions when most of the fine-grained sediment, along with the part of fine 322 sand, becomes the main component of the suspended load. Taking into account these 323 components, we have covered all the material on the vertical profile and can reasonably 324 conclude on the representativeness of the material collected by TIMS. The cross-section 325 sampling was not performed; it is assumed that fine-grained material is rather 326 homogeneously distributed across the horizontal profile (Walling et al., 2000; Perks et al., 327 2014). The geochemical normalization of materials was performed using enrichment factor 328 (EF); all elements in a single-point SPM (average concentrations of all samples except during 329 high water level) and fine-grained sediment were doubly-normalized to TIMS sample using a 330 formula:

331  $EF(E) = (E/AI)_{SPM/sediment} / (E/AI)_{TIMS}$ 

[1]

where E is the element of interest normalized to insoluble element (AI) to minimize dilution
caused by quartz, carbonates, or organic matter. Here, AI is chosen as the best reference
non-mobile element, which is minimally affected by hydrodynamic sorting (Garzanti et al.,
2013). EF > 1 indicates enrichment of element compared to TIMS, while EF < 1 indicates a</li>
depletion.

337 **3.5.1**. The Sava River

338 The chemical composition of suspended sediment usually changes throughout the 339 hydrological cycles (Viers et al., 2008). The temporal variability of flow can cause sorting 340 processes that are responsible for the geochemical differentiation of suspended material 341 based on hydrodynamic properties. During the first autumn sampling campaign (Fig. 6A) the 342 normalized elemental composition of sediment shows enrichment pattern for Co, Sn and Pb. 343 Even though all analyzed materials are under the anthropogenic impact (Milačič et al., 2017), 344 these anomalies suggest that the sediment is more influenced by anthropogenic 345 contamination than the material collected by TIMS. In addition, slight enrichment of Cs, Be, 346 LREE, Th, U, Mo, Fe, Ni, Cu, Zn, Ga and Tl in sediment compared to TIMS is related to

347 elements mainly hosted in micas or associated with clay, Fe-Mn oxyhydroxides and organic 348 matter, which indicates a minor loss of the suspended particles from the sampler. Contrary to 349 sediment, an average of elements in all single-point SPM samples (average of 5 samples) 350 shows lower concentrations of Na, Rb, Mg, Ca, Sr, Ba, Sc, Y, Eu, Lu, Ti, V, Nb, Cr, W, Co, 351 Ni, Ge, Sn, and higher concentrations of K, Be, Th, Mn, Zn and Pb in relation to TIMS. The 352 depletion of the first group of elements is caused by their tendency to accumulate in deep 353 suspended load (Garzanti et al., 2011; Bouchez et al., 2011; Wu et al., 2013) and due to their 354 hydrodynamic properties could be captured in the sampler at high water level when the part 355 of bed load is also taken in suspension. It is assumed that Ca, Mg and Sr are mainly hosted 356 in carbonates, Na, Sr and Eu in tectosilicates, while Lu, Ti, Nb, Cr, Ni, Ge and Sn can be 357 hosted partly in ultradense minerals. Higher enrichment of Mn in single-point SPM is a result 358 of its presence in clay minerals and oxyhydroxides, which dominate in a shallow suspended 359 load and are likely to escape from sampler. Nevertheless, the normalized pattern of average 360 SPM assumes reasonably good representativeness of material collected by TIMS.

361 **Figure 6.** 

362 The second autumn sampling campaign (Fig. 6B) is characterized by more extreme 363 hydrological conditions (Fig. 2). In that period, single-point SPM was sampled mostly at 364 medium discharges. When compared to TIMS, sediment displays a similar pattern as 365 observed in the previous period, i.e. minor enrichment of mostly geogenic (Na, Mg, REE and 366 Th) and some of anthropogenic elements (Ni, Zn, Cd, Sn, Pb and As). With respect to most 367 of the naturally-derived elements (from Li to Nb, W, Fe, Ga and Ge), average SPM values 368 show good similarity to TIMS. Moreover, the SPM at high water level does not show 369 significant enrichment which suggests that most of the material is captured inside the 370 sampler. Depletions of Mg, Ca, Sr could be explained by more siliciclastic supply in 371 conditions of high discharges, when the Sava River brings more material from the upper 372 reaches. Opposite variations of Mo, Mn, Cu, Zn and Cd during high and low water levels are 373 probably a consequence of their different source nature. During high discharges, the SPM is 374 mainly detrital in origin and metal-poor (Ollivier et al., 2011). Contrary, low water conditions

375 are characterized by metal-rich fraction which tends to adsorb onto clays, oxyhydroxides and 376 organic matter (Horowitz and Elrick, 1987; Baran et al., 2019). Based on that, we can 377 assume that TIMS sample represents averaged out SPM composition fairly well. 378 In the winter sampling period (Fig. 6C), single-point SPM samples were collected during low 379 discharges (Fig. 2). In these conditions, for most of the period, the lowest SPM concentration 380 was observed. Here geochemical differences between different water stages are more 381 pronounced. High enrichment patterns for Mo, Mn, Ni, Cu, Zn, Cd, Sn, As, Sb and Bi at low 382 water stage may emphasize prominent instant anthropogenic input and preferable adsorption 383 onto clays/oxyhydroxides. During high discharges, siliciclastic and detrital influence is more 384 visible from the depletion of Mg, Ca and Sr in comparison to the previous sampling period. 385 Grain-size and mineralogical data suggest a good representativeness of material captured in 386 the sampler. This is because both analyses showed a higher amount of clay fraction and 387 phyllosilicates (Fig. 3 and 4) compared to sediment. Also, this is supported by no enrichment 388 patterns in average SPM for most of the geogenic elements (Li. Na, K, Rb, Cs, Be, Ba, Sc, 389 REE, Th, Ti, V, Nb, Fe, Ga and Ge). However, geochemical normalization of high water SPM 390 and sediment samples showed slight enrichments of REE and Th in relation to TIMS. 391 According to Garzanti et al., (2011) these elements can be enriched in both, shallow and 392 deeper suspended load, mainly as a result of provenance effect, concentration in ultradense 393 minerals, phyllosilicates and organic matter. Herein, we assumed that enrichment in 394 SPM/sediment sample could reflect instant provenance signal and scavenging of Th and 395 particularly MREE (Sm, Eu, Gd and Tb) onto Mn/Fe-oxyhydroxides and clays (Quinn et al., 396 2006). 397 The spring sampling campaign (Fig. 6D) is characterized by calm hydrological conditions, 398 disturbed during the last three days of sampling (Fig. 2). In this period, the highest discharge 399 and SPM concentration were observed. The results obtained for sample from deep-

400 positioned TIMS corroborated previously stated inferences about hydraulic behavior of

401 elements; enrichment/accumulation of Na, Mg, Ca, Sr, Ba, Ti, Nb, Cr, Ge and Sn in the bed

402 load. Other enriched values of REE, V, Fe, Co, Ni and Cu suggest that these elements are

403 not solely dominant in a shallow load, but also partly transported in a deeper load hosted in 404 heavy minerals (Garzanti et al., 2011). Based on slight enrichment of mostly geogenic 405 elements (K, Rb, REE, U) in sediment and SPM at high water stage, together with lower 406 mean particle-size observed in sediment sample, a minor loss of finest particles from TIMS is 407 presumed. Similarly to previous sampling period, consecutive enrichments of Mo, Mn, Ni, Cu, 408 Zn, Cd, Sn and Pb are observed. This may suggest that comparing potentially anthropogenic 409 elements in different materials should be taken with care since this group of elements 410 represents fraction more sensitive to physico-chemical changes (pH, redox condition, 411 temperature, electric conductivity, a form of elements, etc.) in water. Therefore, in evaluating 412 time-integrated sampling in human-impacted rivers, focus should be put on naturally-derived 413 elements that are more stable and associated with the residual fraction (Aguilar-Hinojosa et 414 al., 2016; Baran et al., 2019). 415 As discussed earlier, the summer sampling period (Fig. 6E) was characterized by low 416 discharges and unusually high concentration of Ca in shallow-positioned sampler. 417 Considering that Ca tends to accumulate in coarser sediment fraction, that may imply the 418 origin of Ca different than only detrital (Chen et al., 2014). An explanation could be intense 419 biological production at low water stage during which the elements are accumulated by the 420 long-chain organic acids (Rogerson et al., 2008). In addition, slow-moving water conditions in 421 the sampler promoted biogenic calcite precipitation (Olivier et al., 2011), which resulted in 422 unusually high concentration of calcium. These processes invoked the disturbance of the 423 chemical composition of collected material which made the conclusion about sampler's 424 representativeness more difficult to reach. This is most apparent in higher enrichment of all 425 elements (except Ca and Mn) in the sediment compared to TIMS – that would imply the loss 426 of fine material from the sampler which is in contradiction with the PSD results. Higher clay 427 fraction content was determined in shallow TIMS after H<sub>2</sub>O<sub>2</sub> treatment than in sediment. 428 which suggests good effectiveness of TIMS in these low water conditions. However, high Ca

429 content and the evidence of flocculation may mask the real characteristics of the collected

430 material. Therefore, an estimation of the chemical composition in the summer sampling
431 campaign should be interpreted with caution.

For the sampling location Radovljica (the Sava River headwaters), a smaller range of
elements was analyzed (Fig. 7A). The results of single-point SPM have emphasized enriched

- 434 patterns of mostly anthropogenic (Ni, Cu, Zn, Cd and Pb) and naturally derived (Rb, V and
- 435 Mn) elements. As aforementioned, these disparities are related to organic- and metal-rich
- 436 fraction abundant at low water levels. By observing grain-size distribution of the sediment
- 437 samples, TIMS could be evaluated as quite effective. Higher Cr values may reflect

438 anthropogenic impact (Milačić et al., 2017) or appearance of mafic minerals that concentrate

- 439 at the riverbed (Hinterlechner-Ravnik and Pleničar, 1967; Garzanti et al., 2011).
- 440 **Figure 7.**

# 441 3.5.2. Tributaries

The results of normalized diagrams of the Ljubljanica, Savinja and Krapina rivers have shown somewhat different EFs compared to the Sava River (Fig. 7B, C and D). These tributaries have a dissimilar morphology and shallower channels, which promotes coarser particles to be collected by TIMS. This can also explain high quartz content determined in TIMS samples. Discrepancies found between single-point SPM and sediments reflect the influence of hydrodynamic sorting on geochemical and mineralogical composition. Because singlepoint SPM was collected only during low water stage, most of the particles that contain

449 geogenic elements are likely to be captured by TIMS.

TIMSs positioned in the Ljubljanica and Savinja rivers (Fig. 7B and C) were characterized 450 451 with higher concentration of Be and W, particularly in the summer period. The most probable 452 explanation would be their interaction with organic matter (Tuna et al., 2012; Boschi and 453 Willenbring, 2016) or the adsorption on the secondary Fe-Mn oxyhydroxides (Armiento et al., 454 2013; Bauer et al., 2017), rather than the selective entrainment of heavy minerals (Garzanti 455 et al., 2010). As mentioned earlier, long-time sampling inside TIMS can probably cause a 456 change of physico-chemical conditions at the sediment-water interface. The reductive 457 dissolution of oxyhydroxide particles below this layer releases soluble Mn and induces its

458 upward moving into the water column but subsequently trapping it back as Mn oxides when 459 oxygenated water is encountered (Calvert and Pedersen, 1996; Tribovillard et al., 2006). 460 Inside the TIMS, this precipitation resulted in strong binding of Be and W with Mn, and 461 consequently their prominent depletions. This process could also be responsible for depletion 462 patterns of Mn, Pb, Zn, As and Sb observed in sediment samples in comparison to TIMS of 463 the Krapina River (Fig. 7D). As in the case of the Sava River, these variations can reflect 464 their sensitivity to physico-chemical changes at the boundary between sediment and the 465 overlying water (Baran et al., 2019). The higher EFs found in sediments for the most of 466 detrital elements (Cs, Ba, Sc, REE, Th, U, Ga) indicate that their main carriers are not easily 467 retained at high flow rate. This pertains particularly to the Savinja (Fig. 7C), as the 468 hydrologically most demanding river, in which higher loss of particles is expected. Since 469 Ljubljanica and Krapina are typical lowland rivers with lower discharge and lower enrichment 470 patterns, more efficient retention of particles in TIMS is likely to occur.

### **471 4. Conclusions**

In this study, we have combined hydrological observations, granulometric, mineralogical and
geochemical data of the SPM in a medium-size river to characterize element behavior and
representativeness of material collected by time-integrated mass-flux sampler (TIMS). To
evaluate TIMS in real-environment conditions, we have compared different sampled
materials (SPM collected by TIMS, single-point SPM samples and fine-grained sediment (<</li>
63 μm)). The main findings are as follows:
The flocculation process in the river channel induces coarser particle size in the single-

point SPM and the SPM collected by TIMS compared to fine-grained sediments (< 63µm),

480 which otherwise represent their suitable representative. After organic matter removal, a

481 notable increase in clay content and lower mean grain size were observed.

482 2. The mineral composition of all samples is dominated by quartz, carbonates, phyllosilicates

483 and feldspars which are mainly detrital in origin. However, a high content of calcite in

484 sediment (55%) and shallow TIMS (69%) determined in the summer period can be the result

of biologically instigated carbonate precipitation, supported by lower water table and almost
stagnant water conditions.

487 3. The geochemistry of analyzed materials is mostly influenced by the hydrodynamic sorting 488 and provenance; different compositions during different hydrological regimes are found. In 489 calm hydrological conditions, surface load shows enrichment patterns of partly anthropogenic 490 elements (Mo, Mn, Cu, Zn, Cd, As, Sb and Bi) adsorbed onto clay minerals, oxyhydroxides 491 and organic matter. The differentiation of suspended material is not observed at high water 492 stages when more detrital material is supplied, and part of a bed load is re-suspended. 493 4. During low and medium water table, samplers at the Sava River proved to be a reasonably 494 good means of collecting representative time-integrated suspended material. However, 495 different chemical composition of shallow and deep-positioned sampler, induced by 496 hydrodynamic sorting, could be recognized. Somewhat problematic conditions can occur at 497 high flow rate due to partial loss of the clay fraction and variation of elemental composition, 498 particularly in human-impacted rivers. Therefore, using anthropogenic elements, mostly 499 bound to that fraction, to compare materials sampled in different periods, requires additional 500 caution. Moreover, changes in redox conditions and high biological activity in a summer 501 period can invoke additional modifications of the chemical composition of the material and 502 consequently distort the conclusion about the representativeness of the SPM. Hence, in 503 order to interpret geochemical and mineralogical data (time-integrated) in sediment source 504 modeling or assessing river sediment quality, these factors should be considered.

- 505 Acknowledgements
- 506 This work was supported by the Croatian Science Foundation under the project 7555
- 507 (TRACESS) and bilateral Croatian-Slovenian MSES project. The authors are grateful to Dr.
- 508 Ivo Lučić for his support in the field during the sampling campaigns as well as Mr. Roman
- 509 Trček's help on finding an appropriate location for TIMS positioning in the Savinja River.
- 510 **5. References**
- 511 Aguilar-Hinojosa, Y., Meza-Figueroa, D., Villalba-Atondo, A. I., Encinas-Romero, M. A.,
- 512 Valenzuela-García, J. L., Gómez-Álvarez, A., 2016. Mobility and bioavailability of metals in

- 513 stream sediments impacted by mining activities: the Jaralito and the Mexicana in Sonora,
- 514 Mexico. Water Air Soil Poll 227, 345.
- 515 Armiento, G., Bellatreccia, F., Cremisini, C., Ventura, G.D., Nardi, E., Pacifico, R., 2013.
- 516 Beryllium natural background concentration and mobility: a reappraisal examining the case of
- 517 high Be-bearing pyroclastic rocks. Environ Monit Assess 185, 559–572.
- 518 Audry, S., Schäfer, J., Blanc, G., Bossy, C., Lavaux, G., 2004. Anthropogenic components of
- 519 heavy metals (Cd, Zn, Cu, Pb) budgets in the Lot-Garonne fluvial system (France). Appl
- 520 Geochem 19(5), 769–786.
- 521 Baran, A., Mierzwa-Hersztek, M., Gondek, K., Tarnawski, M., Szara, M., Gorczyca, O.,
- 522 Koniarz, T., 2019. The influence of the quantity and quality of sediment organic matter on the
- 523 potential mobility and toxicity of trace elements in bottom sediment. Environ Geochem Hlth
- 524 41, 2893–2910.
- 525 Bauer, S., Blomqvist, S., Ingri, J., 2017. Distribution of dissolved and particulate
- 526 molybdenum, vanadium, and tungsten in the Baltic Sea. Mar Chem 196, 135–147.
- 527 Boschi, V., Willenbring, J.K., 2016. Beryllium desorption from minerals and organic ligands 528 over time. Chem Geol 439, 52–58.
- 529 Bouchez, J., Gaillardet, J., France-Lanord, C., Maurice, L., Dutra-Maia, P., 2011. Grain size
- 530 control of river suspended sediment geochemistry: clues from Amazon River depth profiles.
- 531 Geochem Geophy Geosy 12:Q03008.
- 532 Calvert, S.E., Pedersen, T.F., 1996. Sedimentary geochemistry of manganese: implications
- 533 for the environment of formation of manganiferous black shales. Econ Geol 91, 36–47.
- 534 Cerar, S. Urbanc, J., 2013. Carbonate Chemistry and Isotope Characteristics of Groundwater
- 535 of Ljubljansko Polje and Ljubljansko Barje Aquifers in Slovenia. Sci World J 2013, 1–11.

- 536 Chen, J.B., Gaillardet, J., Bouchez, J., Louvat, P., Wang, Y.N., 2014. Anthropophile elements
  537 in river sediments: overview from the Seine River, France. Geochem Geophy Geosy
  538 15:4526–4546.
- 539 Droppo, I.G., di Cenzo, P., Parrott, J., Power, J., 2019. The alberta oil sands eroded
- 540 bitumen/sediment transitional journey: influence on sediment transport dynamics, pah
- signatures and toxicological effect. Sci Total Environ. 677, 718–731.
- 542 Droppo, I.G., Ongley, E.D., 1994. Flocculation of suspended sediment in rivers of south-543 eastern Canada. Water. Res 28, 1799–1809.
- 544 Fiket, Ž., Mikac, N., Kniewald, G., 2017. Mass fractions of forty-six major and trace elements,
- 545 including rare earth elements, in sediment and soil reference materials used in environmental
- 546 studies. Geostand Geoanal Res 41, 123–135.
- 547 Galović, L., Peh, Z., 2014. Eolian contribution to geochemical and mineralogical
- 548 characteristics of some soil types in Medvednica Mountain, Croatia. Catena 117, 145–156.
- 549 Garçon, M., Chauvel, C., France-Lanord, C., Limonta, M., Garzanti, E., 2014. Which minerals
- 550 control the Nd–Hf–Sr–Pb isotopic compositions of river sediments? Chem Geol 364, 42–55.
- 551 Garzanti, E., Andò, S., 2019. Heavy Minerals for Junior Woodchucks. Minerals 9, 148.
- 552 https://doi.org/10.3390/min9030148.
- 553 Garzanti, E., Ando, S., France-Lanord, C., Censi, P., Vignola, P., Galy, V., Lupker, M., 2011.
- 554 Mineralogical and chemical variability of fluvial sediments 2. Suspended-load silt (Ganga-
- 555 Brahmaputra, Bangladesh). Earth Planet Sci Lett 302, 107–120.
- 556 Garzanti, E., Andò, S., France-Lanord, C., Vezzoli, G., Galy, V., Najman, Y., 2010.
- 557 Mineralogical and chemical variability of fluvial sediments. 1. Bedload sand (Ganga-
- 558 Brahmaputra, Bangladesh). Earth Planet Sci Lett 299, 368–381.
- 559 Garzanti, E., Andò, S., Vezzoli, G., 2009. Grain-size dependence of sediment composition
- and environmental bias in provenance studies. Earth Planet Sci Lett 277(3-4), 422–432.

- 561 Garzanti, E., Padoan, M., Andò, S., Resentini, A., Vezzoli, G., Lustrino, M., 2013. Weathering
- 562 and relative durability of detrital minerals in equatorial climate: sand petrology and
- 563 geochemistry in the East African Rift. J Geol 121, 547–580.
- 564 Gippel, C.J., 1995. Potential of turbidity monitoring for measuring the transport of suspended 565 solids in streams. Hydrol Process 9(1), 83–97.
- Goharrokhi, M., Pahlavan, M., Lobb, D.A., Owens, P.N., Clark, S.P., 2019. Assessing issues
  associated with a time-integrated fluvial fine sediment sampler. Hydrol Process 33, 15, 2048–
  2056.
- 569 Gregory, J., 2006. Particles in Water: Properties and Processes. IWA Publishing/CRC.
- 570 Press, London, U.K.
- 571 Hinterlechner-Ravnik, A., Pleničar, M., 1967. Smrekovški andezit in njegov tuf (The
- 572 Smrekovec and esite and its tuff in Slovenian). Geologija 10, 219–237.
- 573 Horowitz, A.J., 1991. A Primer in Sediment-trace Element Chemistry. Lewis Publishers,
- 574 Chelsea, MA, USA.
- 575 Horowitz, A.J., Elrick, K.A., 1987. The relation of stream sediment surface area, grain size
- and composition to trace element chemistry. Appl Geochem 2, (4), 437–451.
- 577 Kahle, M., Kleber, M., Jahn, R., 2002. Review of XRD-based quantitative analyses of clay
- 578 minerals in soils: the suitability of mineral intensity factors. Geoderma 109, 191–205.
- 579 Kobold, M., Sušelj, K., 2005. Precipitation forecasts and their uncertainty as input into
- 580 hydrological models. Hydrol Earth Syst Sci 9, 322–332.
- 581 Kynčlová, P., Hron, K., Filzmoser, P., 2017. Correlation between compositional parts based 582 on symmetric balances. Math Geosci 49, 777–796.
- 583 Lerouge, C., David, K., Claret, F., Debure, M., Grangeon, S., Madé, B., Montavon, G.,
- 584 Tournassata, C., 2017. Role of carbonate minerals in the distribution of trace elements in
- 585 marine clay formations. Procedia Earth Planet Sci 17, 798–801.

- 586 Lučić, M., Jurina, I., Ščančar, J. Mikac, N., Vdović, N., 2019. Sedimentological and
- 587 geochemical characterization of river suspended particulate matter (SPM) sampled by time-
- integrated mass flux sampler (TIMS) in the Sava River (Croatia). J Soils Sediments 19, 989–
  1004.
- 590 Martínez-Carreras, N., Krein, A., Gallart, F., Iffly, J.F., Hissler, C., Pfister, L., Hoffmann, L.,
- 591 Owens, P.N., 2012. The Influence of Sediment Sources and Hydrologic Events on the
- 592 Nutrient and Metal Content of Fine-Grained Sediments (Attert River Basin, Luxembourg).
- 593 Water Air Soil Poll 223, 5685–5705.
- 594 Marttila, H., Saarinen, T., Celebi, A., Kløve, B., 2013. Transport of particle-associated
- elements in two agriculture-dominated boreal river systems. Sci Total Environ 461-462:693–
  705.
- 597 McDonald, D.M., Lamoureux, S.F., Warburton, J., 2010. Assessment of a time-integrated 598 fluvial suspended sediment sampler in a high arctic setting. Geogr Ann A, 92A:225–235.
- Milačič, R., Zuliani, T., Vidmar, J., Oprčkal, P., Ščančar, J., 2017. Potentially toxic elements
  in water and sediments of the Sava River under extreme flow events. Sci Total Environ 605–
  606, 894–905.
- Moore, D., Reynolds, R., 1997. X-Ray-Diffraction and the identification and analysis of clay
   minerals. Oxford University Press, New York.
- Morehead, M.D., Syvitski, J.P., Hutton, E.W.H., Peckham, S.D., 2003. Modeling the temporal
  variability in the flux of sediment from ungauged river basins. Global Planet Change 39(1-2),
  95–110.
- 607 Ollivier, P., Radakovitch, O., Hamelin, B., 2011. Major and trace partition and fluxes in the
  608 Rhône River. Chem Geol 285, 15–31.
- 609 Pawlowsky-Glahn, V., Egozcue, J. J., Tolosana-Delgado, R., 2015. Modeling and Analysis of
- 610 Compositional Data. London: John Wiley & Sons.

611 Perks, M.T., Warburton, J., Bracken, L., 2014. Critical assessment and validation of a time-

612 integrating fluvial suspended sediment sampler. Hydrol Process 28(17), 4795–4807.

613 Phillips, J.M., Russel, M.A., Walling, D.E., 2000. Time-integrated sampling of fluvial

suspended sediment: a simple methodology for small catchments. Hydrol Process 14, 2589–2602.

616 Placer, L., 2008. Principles of the tectonic subdivision of Slovenia. Geologija, 51/2, 205–217.

617 Quinn, K.A., Byrne, R.H., Schijf, J., 2006. Sorption of yttrium and rare earth elements by

amorphous ferric hydroxide: influence of pH and ionic strength. Mar Chem 99, 128–150.

619 R core team, 2017. R: a language and environment for statistical computing. R Foundation

620 for statistical computing, Vienna, Austria URL http://www.R-project.org.

Rambeau, C.M.C., Baize, D., Saby, N.P.A., Matera, V., Adatte, T., Foellmi, K.B., 2010. High
Cadmium concentrations in Jurassic limestone as the cause for elevated cadmium levels in
deriving soils: a case study in Lower Burgundy, France. Environ Earth Sci 61, 1573–585.

Reimann, C., Filzmoser, P., Fabian, K., Hron, K., Birke, M., Demetriades, A., 2012. The

625 concept of compositional data analysis in practice – total major element concentrations in 626 agricultural and grazing land soils of Europe. Sci Total Environ 426, 196–210.

Reimann, C., Filzmoser, P., Hron, K., Kynčlová, P., Garrett, R.G., 2017. A new method for
correlation analysis of compositional (environmental) data – a worked example Sci Total
Environ 607–608, 965–971.

Rogerson, M., Pedley, H. M., Wadhawan, J. D., Middleton, R., 2008. New Insights into
Biological Influence on the Geochemistry of Freshwater Carbonate Deposits. Geochim
Cosmochim Ac 72, 4976–4987.

Russell, M.A., Walling, D.E., Hodgkinson, R.A., 2000. Appraisal of a simple sampling device
for collecting time-integrated fluvial suspended sediment samples. International Association
of Hydrological Sciences 263.

Salminen, R., Batista, M.J., Bidovec, M., Demetriades, A., De Vivo, B., De Vos, W., Duris, M.,
Gilucis, A., Gregorauskiene, V., Halamić, J., et al., 2005. FOREGS Geochemical Atlas of
Europe, Part 1: Background Information, Methodology and Maps. Geological Survey of
Finland.

Schindler Wildhaber, Y., Michel, C., Burkhardt-Holm, P., Bänninger, D., Alewell, C., 2012.
Measurement of spatial and temporal fine sediment dynamics in a small river. Hydrol Earth
Syst Sc 16, 1501–1515.

- 643 Schultz, L.G., 1964. Quantitative interpretation of mineralogical composition from X-Ray and
- 644 chemical data for the Pierre Shale. U. S. Geological Survey, Professional Paper 391-C.
- 645 Singh, P., 2009. Major, trace and REE geochemistry of the Ganga River sediments:
- 646 Influence of provenance and sedimentary processes. Chem Geol 266(3-4), 242–255.
- 647 Smith, B.T., Owens, P.N., 2014. Flume- and field-based evaluation of a time-integrated
- suspended sediment sampler for the analysis of sediment properties. Earth Surf Proc Land39, 1197–1207.
- 50 Šikić, K., Basch, O, Šimunić, A., 1979. Geological Map of SFRJ in scale 1:100000.
- 651 Explanatory booklet to sheet Zagreb.- Federal Geological Survey, Beograd.
- Tribovillard, N., Algeo, T., Lyons, T.W., Riboulleau, A., 2006. Trace metals as paleoredox
- and paleoproductivity proxies: an update. Chem Geol 232, 12–32.
- Tuna, G.S., Braida, W., Ogundipe, A., Strickland, D., 2012. Assessing tungsten transport in
- thevadose zone: From dissolution studies to soil columns. Chemosphere 86, 1001–1007.
- Viers, J., Dupré, B., Gaillardet, J., 2009. Chemical composition of suspended sediments in
- 657 World Rivers: New insights from a new database. Sci Total Environ 407, 853–868.
- Viers, J., Roddaz, M., Filizola, N., Guyot, J.L., Sondag, F., Brunet, P., Zouiten, C.,
- 659 Boucayrand, C., Martin, F., Boaventura, G.R., 2008. Seasonal and provenance controls on

- 660 Nd-Sr isotopic compositions of Amazon rivers suspended sediments and implications for Nd
- and Sr fluxes exported to the Atlantic Ocean, Earth Planet Sci Lett 274, 511–523.
- 662 Walling, D.E., Owens, P.N., Waterfall, B.D., Leeks, G.J., Wass, P. D., 2000. The particle size
- 663 characteristics of fluvial suspended sediment in the Humber and Tweed catchments, UK. Sci
- 664 Total Environ, 251-252, 205–222.
- 665 Woodward, J., Walling, D.E., 2007. Composite suspended sediment particles in river
- systems: their incidence, dynamics and physical characteristics. Hydrol Process 21, 3601–3614.
- 668 Wu, W., Zheng, H., Xu, S., Yang, J., Liu, W., 2013. Trace element geochemistry of riverbed
- and suspended sediments in the upper Yangtze River. J Geochem Explor 124, 67–78.

671	Figure caption
672	
673	Figure 1. Map of sampling sites (map sourced from <u>https://maps-for-free.com/</u> ).
674	
675	Figure 2. Water discharge and suspended particulate matter (SPM) concentrations at the
676	Zagreb sampling site during five sampling campaigns (October and November 2016; February,
677	April and July 2017). Data provided from Meteorological and Hydrological Service.
678	Colored circles represent the frequency of collecting single-point SPM and sediment samples.
679	
680	Figure 3. Particle size distribution curves of investigated samples from the Sava River, Zagreb;
681	A) all samples and B) samples presented as average values.
682	
683	Figure 4. Mineral composition of analyzed samples.
684	
685	Figure 5. Heat-maps of correlation coefficients based on symmetric coordinates for different
686	types of samples; A) TIMS, B) single-point SPM and C) fine-grained sediments. Samples from
687	all locations are included.
688	
689	Figure 6. Doubly-normalized chemical composition of fine-grained sediment and SPM from
690	the Sava River, Zagreb. The elements are arranged following the periodic table groups.
691	
692	Figure 7. Doubly-normalized chemical composition of fine-grained sediment and single-point
693	SPM from the Sava River-Radovljica and tributaries. SPM average refers to at least four single-
694	point SPM samples. The samples are normalized to TIMS. The elements are arranged
695	following the periodic table groups.
696	
697	
698	
699	
700	